COMPARATIVE ANALYSIS OF THE SOIL EROSION IN HILL BASINS (SĂSĂUŞ AND MISLEA)

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The present study focuses on soil surface erosion and applying the Universal Soil Loss Equation with GIS techniques of spatial analysis on an area of two river basins. The erosion and the processes associated with it are studied with the help of digital terrain data and the USLE and RUSLE models are successfully applied within the area. Soil surface erosion occurs when detachable soils on sufficiently steep slopes are exposed to overland flow and/or the impact of rainfall.

The Universal Soil Loss Equation (USLE) predicts the long term average annual rate of erosion on a field slope based on rainfall pattern, soil type, topography, crop system and management practices but does not however predict the soil loss resulting from gully erosion. Five major factors are used to calculate the soil loss for a given site. Each factor is the numerical estimation of a specific condition that affects the severity of soil erosion at a particular location.

Key words: GIS techniques, soil erosion, Universal Soil Loss Equation, Săsăuş River Basin, Mislea River basin.

Introduction

Addressing and combating soil erosion assessment dates from the third decade of the twentieth century as a reaction against the effects of grubbing and massive deforestation (Iowa State University, 1926). In Europe, only the Austrian and Swedish agriculture has contributed to increase fertility in soils affected by erosion, due to their protective measures, but afterwards entered into decline.

Over time several stages have been revealed in the research models based on physical soil erosion, which allow an adequate and a quantitative estimate, taking into account both erosion and deposition on a regional scale.

During the simple models stage, the obtained results were completed by developing the Universal Soil Loss Equation (USLE), with the contribution of: Cook M. (1936), Zingg M. (1940), Smith R. M. and Whitt (1947), Musgrave W. G. (1947) and Wischmeier W. H. (1955-1958-1972-1976). The development of

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USLE equation was based on a rich database which was statistically processed. Subsequently from USLE several equations were derived, such as: MUSLE, RUSLE etc.

The development stage of empiric mathematical and deterministic models ended up with the elaboration of the following models: RUSLE (1979) that revises the USLE model, ANSWERS (Beasley and others, 1980), EPIC (Williams J. R., 1985), MORGAN (1995), AGNS.

In the stage of deterministic type models, with a physics base, in which the predominance of empirical equations was limited, the following models were noticed: CREAMS (Knisel W. G. and others, 1980 and Foster, 1981), ROSE (1984), LISEM (1984), WEPP (Foster and Lane,1987; Flanagan D., 1991), LANE (1994), EUROSEM (Morgen and others, 1997), EGEM, EROSION-3D, GLEAMS, KINEROS2, MOSES, MWISED, PESERA, SERAE, STREAM, SWAT, WATEM, CAESAR, WILSIM, WATEM, NRCS, etc.

Since the results of the process models WEPP and EUROSEM, weren't always the best compared to USLE or RUSLE, the review and improvement of the existing physical models represent actual lines of interdisciplinary analysis and connecting to experimental data (Morgan and Nearing, 2000).

In Romania, systematic research regarding surface erosion, performed with drainage plots were initiated by Staicu Ir. (1945), followed by Moţoc M. (1956, 1963, 1973, 1979, 1998), Stănescu P. (1979), Ionescu V. (1981), Ene Al. (1987), Savu P. (1980-1985), Ioniță I. (1990-2000), Pujină D. (1991-1998), Popa N. (2000-2009) etc. Along with the standard plot-level experiments (Rădoane N., 1987; Ioniță I., 2000) researches were also carried out on slopes and hydrographical basins, yielding valuable results (Stanescu P., Drăgan Livia, 1970; Moţoc M., Munteanu S., Băloiu V., Stănescu P., Mihai Gh., 1975; R. J. Bally, 1977; Grecu Florina, 1980; Gaşpar., Untaru E., Roman F., Cristescu C., 1982; Băloiu V., Ionescu V., 1986; Munteanu, S.A., Traci, C., Clinciu, I., Lazăr, N., Untaru, E., Gologan, N., 1993; Pujină Liliana, Pujină D., 1991-1998; Grecu Florina, 1996; Clinciu, I., Lazăr, N., 1997; Rădoane, N., 2003) etc.

Important contributions were made by specialists of the National Institute of Meteorology and Hydrology (INMH), especially those related to production and sediment transport in small and large river basins (Diaconu, Stănescu, Roșca, Miță etc.). Deterministic models have been developed (P. Stanciu, Zlate I., 1985) and water balance modeling was performed in experimental pools (M. Adler, 1993).

Studied Areas

The study area focused on Săsăuş and Mislea hydrographic basins (*Fig. 1*). These basins have many similarities in morphology, despite their location in different morpho-structural units. Both basins have approximately the same area (~200km²), same shape, and order 6 on the Horton-Strahler hierarchy system.

Both basins have a dendritic structure, as evidenced by the existence of tributaries that flow the same direction with the main river; the confluence angles usually less than 90° (Cotet P., 1951).

Săsăuş morpho-hydrographic basin is located within the Romanian territory, and lies at the junction of the Hârtibaciu Plateau with the Făgăraş Depression, in the Transylvanian Depression, framed by the geographical coordinates 24°49'23" and 24°32'14" eastern longitude and 45°56'51" and 45°47'54" northern latitude (*Fig. 2*). Săsăuş basin is found in the southern part of Hârtibaciu Plateau, a clearly defined entity and delimited space in the geographical center of Romania. From a morpho-hydrographical point of view, Săsăuş River falls within Pârâul Nou river basin, a direct tributary to Olt River. The total length of the river Săsauş is 27 km, which drains a surface of 85 km² from the total study area of 232,21 km². The basin is bordered in the north, northwest and west by Hârtibaciu river basin, in the east by Cincu river basin and in the south by Olt river basin. In geological terms the basin area overlaps a Neogene sedimentary package belonging to Sarmatian and Badenian, uncemented rocks (sands and gravels) or weakly cemented rocks (friable sandstone, thin horizons of conglomerates, clays and marls).

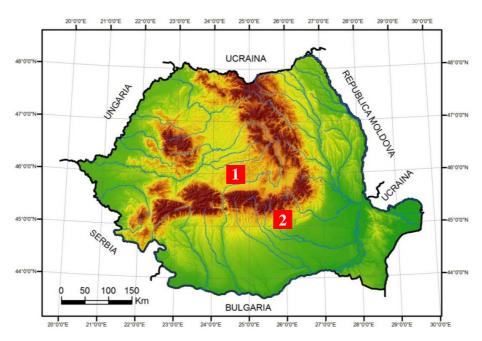


Fig. 1. Geographic location of Săsăuş (1) and Mislea (2) river basins within Romania

Mislea morpho-hydrographic basin is located in the south-eastern part of Romania, at the contact of the Curvature Sub-Carpathians with the Romanian Plain, framed by the geographical coordinates 45°11'25" and 45°03'12" northern latitude, 25°46'19" and 25°59'48" eastern longitude (*Fig. 3*).

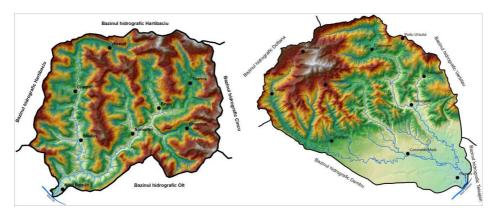


Fig. 2. Săsăuş hydrographic basin, hypsometric map

Fig. 3. Misela hydrographic basin, hypsometric map

The basin has a total area of 175 km² and is a part of Teleajen river basin. It is bordered in the north and east side by Vărbilău river basin, by Doftana basin in the east and Dâmbu basin in the south. From a geological point of view, Mislea basin overlaps the following structural units: Carpathian Molasses, consisting of sandstones, marls, clays, marl-limestone of Mio-Pliocene age and Tarcău nappe consisting of Oligocene and Eocene age formations (sandstones, shales, marls, breccias).

Techniques and Methods

Ability to erosion of an area broadly depends on two important factors: a) heavy erosivity b) soil erodibility. Currently, the research conducted by INCPA Bucharest (simulation models of water balance in soil, soil loss forecasting models for agricultural use by erosion and landslides) and also the Research and Development Center for Combating Soil Erosion in Perieni – Vaslui County, will improve the ROMSEM model (Romania Soil Erosion Model).

 $Table\ 1$ Soil Erosion Models Developed in GIS (after Capolongo, Piccaretta, 2007 – with Additions)

HOLE (II : 10 :1	XX7' 1 ' 1
USLE (Universal Soil	Wischmeier and
Loss Equation)	Smith, 1965
ROMSEM (Romanian Soil	Moţoc and
Erosion Model)	others, 1979
CREAMS (Chemicals,	
Runoff and Erosion	Knisel, 1980
from Agricultural	Killsel, 1960
Management Systems)	
EPIC (Erosion-Productivity	M.111. 1002
Impact Calculator)	Williams, 1985
AGNPS (Agricultural	37 1
Non-Point Source	Young and
Pollution Model)	others, 1989
WEPP (Water Erosion	Nearing and
Prediction Project)	others, 1989
KINEROS (Kinematic	Woolhiser and
Runoff and Erosion Model)	others, 1990
EROSION 3D	Schmidt, 1991, 1996
LISEM (Limburg Soil	De Roo and
Erosion Model)	others, 1996
USPED (Unit Stream Power	Mitasova and
Erosion Deposition)	others, 1996
RUSLE (Revised Universal	Renard and
Soil Loss Equation)	others, 1997
EUROSEM (European Soil	Morgan and
Erosion Model)	others, 1998
PESERA (Pan-European	
Soil Erosion Risk	Gobin and
Assessment Project)	others, 1999
SWAT (Soil and Water	Neitsch and
Assessment Tool)	others, 2001
,	,

For the development of the soil erosion model, topographic maps at 1:25.000 scale (1980 edition), the Romanian Soil Map at scale 1:200.000 and the Corine Land Cover data set from 2006 were used. For the study area these coefficients were calculated at pixel level, for the 20m resolution model. Implementation scheme of the equation is shown in *Fig. 4*.

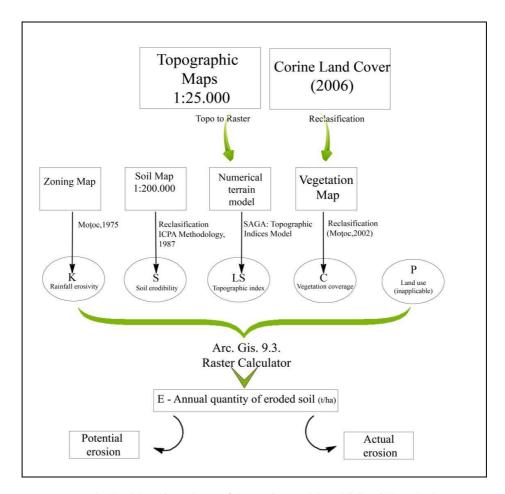


Fig. 4. The elaboration scheme of the erosion model explaining GIS methods (Alexandru, Cătescu, 2012)

Adapting the Universal Soil Loss Equation (USLE) to the specific Romanian conditions, was first attempted by Mircea Moţoc and collaborators in the year 1975. Based on the general equation for assessing potential soil erosion, $Ep = A \cdot L\alpha \cdot I\beta$ (corresponds to the situation when the soil isn't covered with vegetation and anti-soil erosion measures are not taken). Mircea Moţoc developed the actual erosion equation based on the formula presented above, which combines the influence of the slope length (L calculated in meters), slope steepness (I calculated in %) and soil properties in a single parameter:

$$E = K \cdot L^m \cdot I^n \cdot S \cdot C \cdot C_s,$$

where: E = average annual actual erosion rate(t/ha/year);

K = correction coefficient for rainfall erosivity = 100 Eps • Ap / $L^{0.5} ext{ • } I^{1.45}$ (erosion of the standard erosion control plots);

Ap = pluvial aggressiveness = $e^{I15} \cdot H^{0.5}$; L and I = length (m) and slope steepness (%); S = correction coefficient for soil erodibility;

C = correction coefficient for crops effect; Cs = correction coefficient for the effect of erosion control measurements; m = 0.3; $I^n = 1.36 + 0.97i + 0.381i^2$, where

"i" is the average slope.

Rainfall Erosivity (K)

Rainfall erosivity represents the annual sum of the products between the energy of the erosive rainfalls (E) and their maximum 30 minutes intensities:

$$K = \sum_{i=1}^{n} E_i \cdot I30_i$$

where: E_i = the kinetic energy of every rainfall with a duration over 30 minutes during a year (MJ/ha);

 $I30_i$ = the maximum intensity of the 30 minute rainfall (mm/h).

For the studied area, the K factor value was taken from the rainfall erosivity zoning map of Romania (Motoc and others, 1975). This value ranges from 0.12 for the Depression of Transylvania, 0.16 for the Meridional Carpahtians, 0.14 for the Curvature Subcarpathians and 0.13 for the Romanian Plain.

Soil erodibility (S)

This factor was obtained starting from the soil map, taking into account the texture of every soil type, and following the reclassification proposed by ICPA (1987). In the studied areas, soils fall under four classes of erodibility (*Fig. 5*):

• Erodibility class 1 is the most susceptible to erosion and includes: eroded soils, brown eu-mezobasic soils with sandy-clay texture, clay-loamy and clay textures. Within the Mislea river basin, this class is found especially on the slopes of Mislea, Lupăria, Telega and Cosmina valleys and in the Săsăuş river basin on the slopes of Valea Lungă, Valea Ilimbav, Valea Vizina, Gherdeal and Pârâul Nou Valleys.

- Erodibility class 0,80 is characteristic for brown acid soils with sandy-clay texture and its present only within Mislea River basin on the Măceșu Plaiul Rotund interfluve.
- Erodibility class 0,70 is representative for reddish-brown luvic soils, clay-illuvial brown soils, black clinohidromorfic soils, pseudorendzines, brown eu-mezobasic soils, brown luvic soils with clayish texture, sandy-clay and clay textures. This type of class is present in both river basins, especially on the slopes of the following valleys: Cosmina, Doftăneţ, Poiana Trestiei, Gherdeal, Valea Caprelor, Valea lui Trifan, Veseud.
- Erodibility class 0,60 is the class with the lowest tendency to erosion and includes alluvial soils, brown luvic, alluvial protosoils, albic luvisols, with loam-sandy and loamy texture. These soils are developing on the terraces and floodplains of the Mislea, Cosmina, Săsăuş, Şomartin and Pârâul Nou valleys.

The length (L) and slope steepness (I)

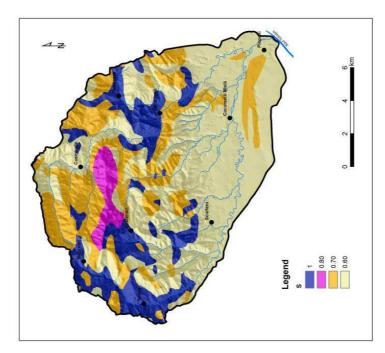
These factors were considered as a single factor (LS) based on the numerical altitude model, to help simplify the analysis in GIS. LS factor causes the influence of the relief on soil erosion and consists in determining the length (L) and slope steepness (S). For the studied areas, this factor was calculated based on the numerical model at a resolution of 20m altitude, applying the Topographic Indices function in SAGA GIS program (Moore, 2003). This factor estimates the highest values up to 400 and correspondes to the areas with the highest tendency to erosion. The highest values of this indicator are found along river thalwegs, ranging between 200 and 400 with the lowest values found on the interfluves, terraces and floodplains (Fig. δ).

The vegetation factor (C)

For the studied areas, the C factor was determined from Corine Land Cover data set (2006) which, after the reclassification, resulted in a number of five classes of vegetation: orchards; pastures, hayfields, grasslands; built area; arable land and forests (*Table 2*).

The corresponding values for factor C, were taken after Moţoc M. (1975), assigned to each class of vegetation and processed in GIS, which led to creating the distribution map of the C factor.

From the analysis of the two maps (Fig. 7), it results that the highest value of 0.50 is assigned to orchards, the value of 0.45 corresponds to arable land with a higher proportion in Mislea basin, the value of 0.30 corresponds to pastures, hayfields and grasslands with a higher proportion found in Săsăuş basin, 0.20 value represents the forests and the 0.02 value is attributed to the built perimeter.



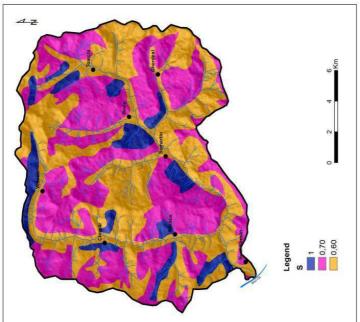


Fig. 5. The distribution of the soil erodibility Factor (S) within Săsăuş and Mislea river basins

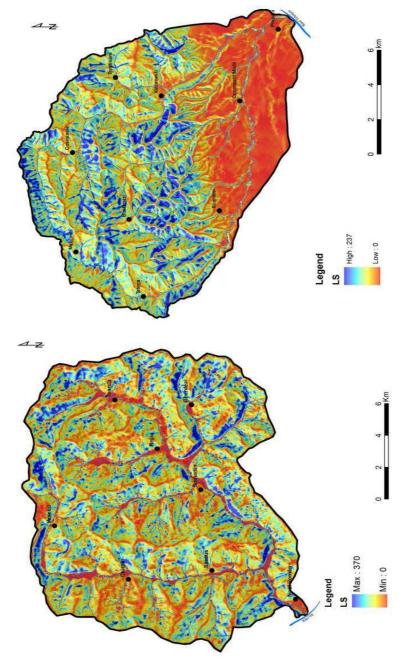


Fig. 6. The distribution of the length and slope steepness Factor (LS) within Sāsāuş and Mislea river basins

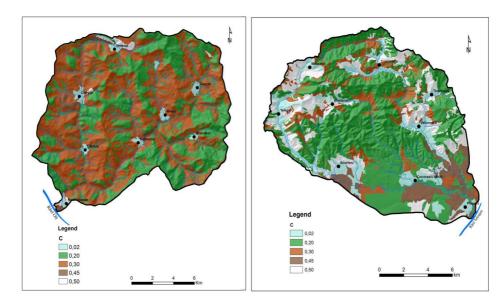


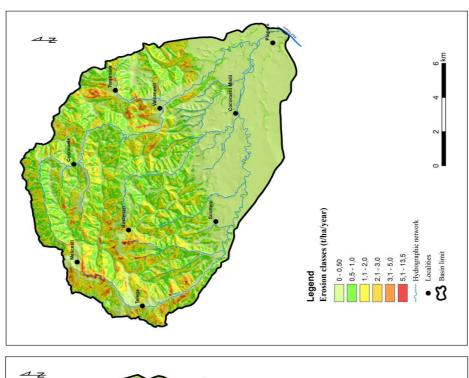
Fig. 7. The distribution of the vegetation coverage Factor (C) within Săsăuş and Mislea River basins

The Correction Coefficient for the Effect of Erosion Control Measurements (P/Cs)

This factor was not considered due to limited applicability and lack of data.

Results and Discussions

Maps of actual and potential erosion (*Fig.* 8 and *Fig.* 9) resulted from the combination of rasters corresponding factors K, LS, S and C, respectively K, LS and S factors. From the analysis of the obtained maps the following observations emerged: the actual erosion model estimates an annual quantity of eroded soil ranging from 0 to 8.4 t/ha/year (Săsăuş Basin), and 0 to 13.5 t/ha/year (Mislea Basin). Most of Săsăuş Basin falls within the ranges of 0-1 t/ha/year, while the share of Mislea Basin erosion is given by the intervals 0-2 t/ha/year. The lowest erosion values (0-1 t/ha/an) can be found particularly in areas covered with forests and along the main river valleys of Săsăuş and Pârâul Nou.



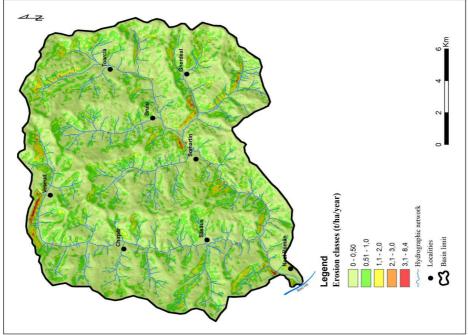
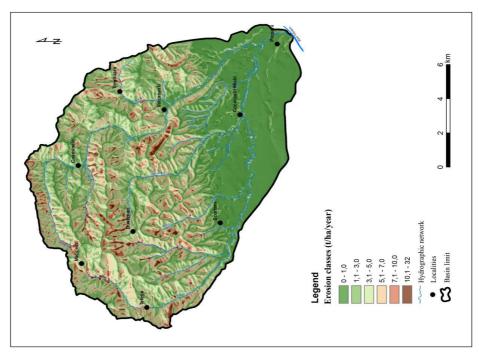
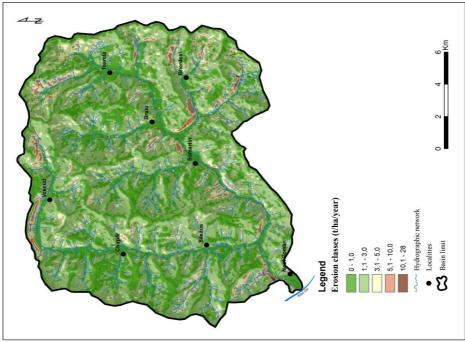


Fig. 8. Actual Erosion within Săsăuş and Mislea River Basins





 $\it Fig.~9$. Potential Erosion within Săsăuş and Mislea River Basins

Regarding Mislea Basin, the lowest values have a significant predominance in its southern region and along Mislea, Seaca Cosmina and Doftăneţu Valleys.

Within both basins the following differences are noticed: the highest values of actual erosion within Mislea Basin correspond to torrential river basins with steep slopes and on the slopes of Cosmina and Telega valleys. Meanwhile, within Săsăuş basin, the highest values are distributed in compact areas in the southern part of the basin, along the Pârâul Nou, Valea Lungă, Gherdeal Valleys and in the northern region of Vizina Valley.

The highest values of the erosion correspond to the surfaces occupied by herbaceous vegetation (pastures, hayfields, grasslands) and orchards, overlapped on beds of clay, sandy-argillaceous texture soils and on very steep slopes. Following a scenario of potential erosion, where the vegetation factor (C) was removed, erosion values have resulted within the ranges of 0-28 t/ha/year for Săsăuş basin and 0-32 t/ha/year for Mislea basin. The lowest values generally correspond to those obtained in the case of actual erosion in both basins.

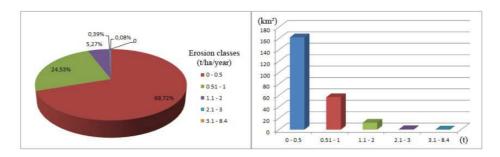


Fig. 10. The Cyclogram and Histogram of the Actual Erosion within Săsăuş River Basin

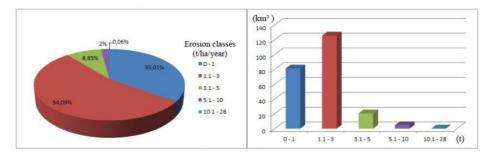


Fig. 11. The Cyclogram and Histogram of the Potential Erosion within Săsăuş River Basin

In the Săsăuş river basin, the highest values are found in areas consistent with those of the actual erosion, obviously with almost a triple amount compared to the initial value (Valea Lungă, Pandea, Gherdeal, Pârâul Nou, Valea Vizina between 3-28 t/ha/year). Mislea hydrographical basin has a higher degree of

potential erosion compared to Săsăuş, with 32 t/ha/year as a maximum value. The highest values can be found in the whole northern half of the basin, focusing on the slopes of Telega, Runcu, Doftăneţu, Poiana Trestia and Lupăria. Considering the classification of ICPA for susceptibility classes it is noticed that following the correlation with the amount of eroded soil, Săsăuş river basin falls within the classes <1 t/ha/year and 1-8 t/ha/year, with a small or sometimes absent susceptibility to erosion.

Regarding Mislea basin, the situation is similar, but due to a higher amount of eroded soil (up to 13.5 t/ha/year), we also take into account a moderate susceptibility class (8-16 t/ha/year). Thus, for both basins one cannot speak of increased erosion, as its maximum values do not exceed 8 and 13.5 t/ha/year, so the development of a susceptibility map is not relevant.

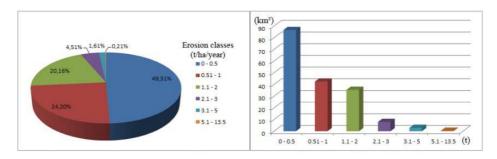


Fig. 12. The Cyclogram and Histogram of the Actual Erosion within Mislea River Basin

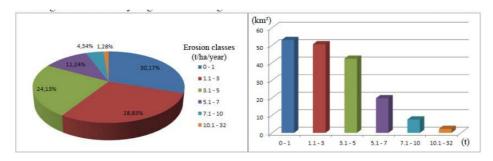


Fig. 13. The Cyclogram and Histogram of the Potential Erosion within Mislea River Basin

Conclusions

The method used was successfully applied, reflecting the reality on the field. This allows the assessment of the degree of surface erosion within the two analyzed basins and the design of protective measures for areas affected by high

erosion. After the analysis of the two basins the average of the surface erosion per year which resulted is of 0.6 t/ha.

This result concludes that the basins under study do not fit within a high risk erosion situation. Taking into account different scenarios for the level of vegetation coverage or land use, we can assess the effects of erosion on the landscape and propose strategies to combat erosion. Through the use of GIS techniques in this study, erosion is highlighted by obtaining maps of actual and potential erosion. Regarding these possibilities, we can carry out monitoring and maintenance on the slopes soil quality, which would allow for appropriate protection measures to be taken.

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